

# Projected land photosynthesis constrained by changes in the seasonal cycle of atmospheric CO<sub>2</sub>

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Uncertainties in the response of vegetation to rising atmospheric CO<sub>2</sub> concentrations<sup>1,2</sup> contribute to the large spread in projections of future climate change<sup>3,4</sup>. Climate–carbon cycle models generally agree that elevated atmospheric CO<sub>2</sub> concentrations will enhance terrestrial gross primary productivity (GPP). However, the magnitude of this CO<sub>2</sub> fertilization effect varies from a 20 per cent to a 60 per cent increase in GPP for a doubling of atmospheric CO<sub>2</sub> concentrations in model studies<sup>5–7</sup>. Here we demonstrate emergent constraints<sup>8–11</sup> on large-scale CO<sub>2</sub> fertilization using observed changes in the amplitude of the atmospheric CO<sub>2</sub> seasonal cycle that are thought to be the result of increasing terrestrial GPP<sup>12–14</sup>. Our comparison of atmospheric CO<sub>2</sub> measurements from Point Barrow in Alaska and Cape Kumukahi in Hawaii with historical simulations of the latest climate–carbon cycle models demonstrates that the increase in the amplitude of the CO<sub>2</sub> seasonal cycle at both measurement sites is consistent with increasing annual mean GPP, driven in part by climate warming, but with differences in CO<sub>2</sub> fertilization controlling the spread among the model trends. As a result, the relationship between the amplitude of the CO<sub>2</sub> seasonal cycle and the magnitude of CO<sub>2</sub> fertilization of GPP is almost linear across the entire ensemble of models. When combined with the observed trends in the seasonal CO<sub>2</sub> amplitude, these relationships lead to consistent emergent constraints on the CO<sub>2</sub> fertilization of GPP. Overall, we estimate a GPP increase of  $37 \pm 9$  per cent for high-latitude ecosystems and  $32 \pm 9$  per cent for extratropical ecosystems under a doubling of atmospheric CO<sub>2</sub> concentrations on the basis of the Point Barrow and Cape Kumukahi records, respectively.

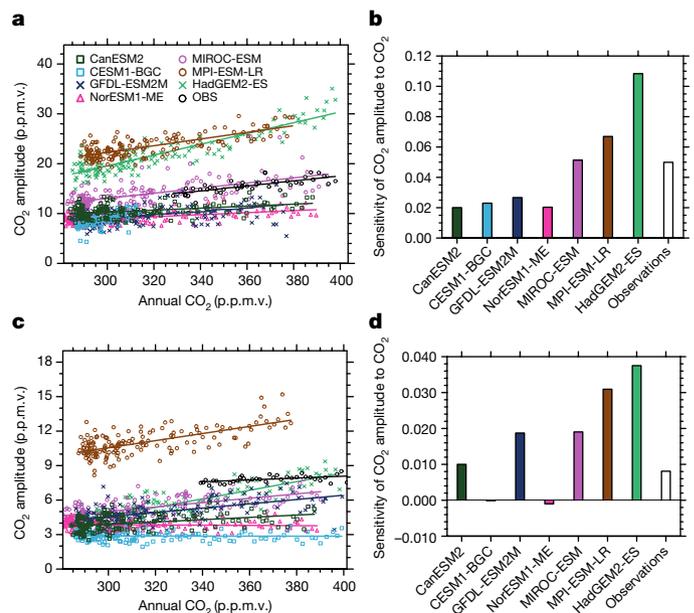
The aim of this study is to reduce the uncertainty in projected large-scale GPP increases, on the basis of the observed trends in the CO<sub>2</sub> amplitude at two measuring sites by applying an emergent constraint<sup>8–11</sup>. This method utilizes common relationships between observables, such as the CO<sub>2</sub> seasonal cycle, and Earth system sensitivities, such as the CO<sub>2</sub> fertilization of the terrestrial carbon sink, considering the full range of responses from an ensemble of complex Earth system models (ESMs).

It has been hypothesized that increasing GPP has been responsible for an observed increase in the amplitude of the CO<sub>2</sub> seasonal cycle at Mauna Loa, Hawaii<sup>14</sup>, but the sensitivity of the seasonal cycle at this high-altitude site to variations in atmospheric circulation has prevented confirmation of this theory<sup>15</sup>. Some recent studies also suggest that variations in the Mauna Loa seasonal cycle are partly due to changing agriculture<sup>12,13</sup>. Here we instead analyse the observed changes in the amplitude of the CO<sub>2</sub> seasonal cycle at Point Barrow, Alaska, a high-latitude station much less affected by mid-latitude agriculture, and at Cape Kumukahi, which is close to Mauna Loa but consists of ground-based measurements that are more directly comparable to the model outputs.

Between 1974 and 2013 the global mean atmospheric CO<sub>2</sub> concentration increased by about 75 p.p.m. by volume (p.p.m.v.) and therefore by about the same amount at Point Barrow (BRW: 71.3° N, 156.6° W),

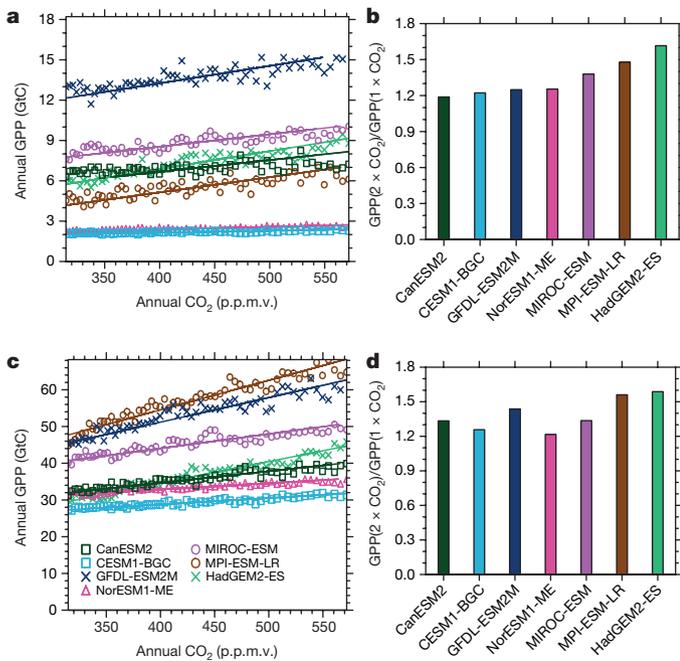
Alaska, and at Cape Kumukahi (KMK: 19.5° N, 155.6° W), Hawaii, (Extended Data Fig. 1). On top of this increasing CO<sub>2</sub> trend, the uptake and release of carbon by the terrestrial biosphere throughout the year causes a seasonal cycle of CO<sub>2</sub>: high concentrations occur in the Northern Hemisphere winter when there is a net release of CO<sub>2</sub> from the land due to the decomposition of organic matter in the soil, and lower values are observed in summer when Northern Hemisphere photosynthesis results in a drawdown of CO<sub>2</sub> (ref. 16). A change in the rate of photosynthesis (for example, due to CO<sub>2</sub> fertilization) or decomposition (due to temperature variability, for instance) will therefore change the amplitude of CO<sub>2</sub> as measured in the atmosphere. In addition, changes in the phase lag between photosynthesis and decomposition, due to the effects of summer drying on photosynthesis or the effects of autumn warming on decomposition<sup>17</sup>, can also change the amplitude of the CO<sub>2</sub> seasonal cycle.

The observed CO<sub>2</sub> amplitude at BRW increased from about 13 p.p.m.v. to 18 p.p.m.v. over the available observational record from 1974 to 2013, and from about 8 p.p.m.v. to 9 p.p.m.v. over the length of



**Figure 1 | Comparison of CO<sub>2</sub> seasonal amplitudes for CMIP5 historical simulations and observations.** **a, c**, Annual mean atmospheric CO<sub>2</sub> versus the amplitudes of the CO<sub>2</sub> seasonal cycle at BRW (**a**) and KMK (**c**) for observations (black) and CMIP5 historical simulations (colours). Markers show the values for the individual years and the lines show the linear best fit for each model and for the observations. **b, d**, Histogram showing the corresponding gradient of the linear correlations for BRW (**b**) and KMK (**d**). Linear trends are derived for the period 1860–2005 from historical simulations for the models, for 1974–2013 for the BRW observations and for 1979–2015 for the KMK observations.

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**Figure 2 | Comparison of simulated annual mean GPP at a doubling of CO<sub>2</sub> in the 1%BGC simulations.** **a, c,** Annual global mean CO<sub>2</sub> increase versus the annual mean GPP in the CMIP5 1%BGC simulations for high-latitude (60° N–90° N) GPP (**a**) and extratropical (30° N–90° N) GPP (**c**). Markers show the values for the individual years and lines show the linear best fit for each model. **b, d,** Histogram showing the relative change in the high-latitude (**b**) and extratropical (**d**) GPP due to a doubling of atmospheric CO<sub>2</sub> (see Methods).

the KMK record (1979–present). To understand and interpret these changes, we make a comparison to ESM simulations that are available via the Coupled Model Intercomparison Project Phase 5 (CMIP5)<sup>18</sup>. We analyse seven ESMs that provided outputs from both a fully coupled historical simulation forced with anthropogenic CO<sub>2</sub> emissions for the period 1850–2005 and a biogeochemically (BGC) coupled simulation that excludes climate change effects and has a prescribed atmospheric CO<sub>2</sub> concentration starting from a preindustrial value for 1850 of around 285 p.p.m.v. and then increasing at 1% per year until quadrupling (hereafter referred to as 1%BGC). The comparison between the historical and 1%BGC runs provides information on the relative influence of CO<sub>2</sub> fertilization and climate change on GPP.

Figure 1a, c compares the observed change in the amplitude of the seasonal CO<sub>2</sub> cycle (black markers) to that simulated by each of the seven CMIP5 models in their historical simulations. For consistency with our subsequent analysis, we have plotted the CO<sub>2</sub> amplitude against the annual mean CO<sub>2</sub> at BRW and KMK. The models simulate the CO<sub>2</sub> amplitudes under the present-day CO<sub>2</sub> concentration (approximately 400 p.p.m.v.) over a range of 12–30 p.p.m.v. at BRW (Fig. 1a) and 3–13 p.p.m.v. at KMK (Fig. 1c). Most models simulate an increase in the CO<sub>2</sub> amplitude over time, but the magnitude of this increase varies considerably from model to model, as shown by the linear regression lines in Fig. 1a, c.

Figure 1b, d compares the gradient of these linear regression lines. The observations (black bar) suggest an increase in the CO<sub>2</sub> amplitude of about 0.05 p.p.m.v. per 1 p.p.m.v. increase in annual mean CO<sub>2</sub> at BRW and 0.008 p.p.m.v. per 1 p.p.m.v. at KMK. The models show a large range in this gradient of about 0.02–0.11 p.p.m.v. per 1 p.p.m.v. at BRW and 0–0.04 p.p.m.v. per 1 p.p.m.v. at KMK. Models with major high-latitude vegetation greening (for example, HadGEM2-ES) give large increases in the CO<sub>2</sub> amplitude, whereas models with strong nitrogen limitations on plant growth (CESM1-BGC, NorESM1-ME)

typically show weaker or even slightly negative trends (Fig. 1d). Overall, weaker trends of around 0.02 p.p.m.v. per 1 p.p.m.v. are favoured by four out of the seven CMIP5 models at BRW (Fig. 1b).

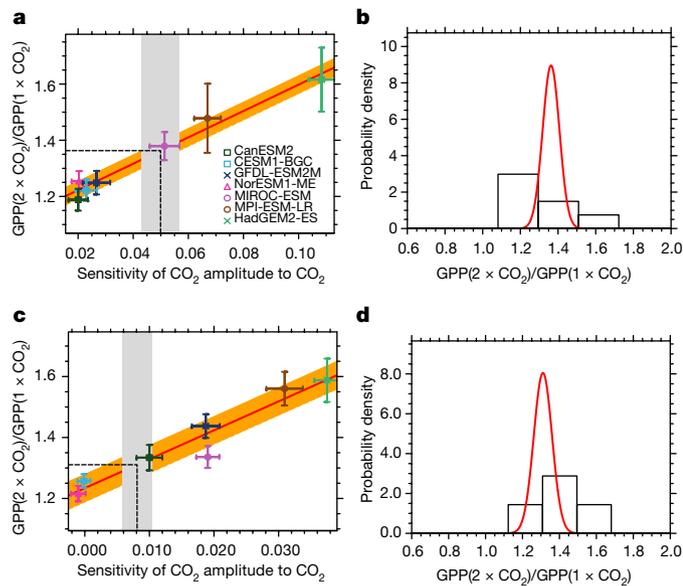
The CO<sub>2</sub> amplitude at BRW is well correlated with annual mean high-latitude GPP in each model, indicating that the dominant cause of the increasing amplitude is increasing GPP (Extended Data Fig. 2). The CMIP5 models agree reasonably well on the gradient of this relationship (0.13–0.22 GtC yr<sup>-1</sup> per p.p.m.v.), which suggests that the observed increase of 5 p.p.m.v. in the CO<sub>2</sub> amplitude at BRW is consistent with an increase of 0.65–1.1 GtC yr<sup>-1</sup> in high-latitude GPP from 1974 to 2013. Changes in the annual mean GPP in the 1%BGC simulations show a slightly smaller rate of increase with CO<sub>2</sub>, implying that high-latitude warming is adding to the increase in CO<sub>2</sub> amplitude that is simulated in the historical runs (Extended Data Fig. 3). However, the simulated overall increase in the CO<sub>2</sub> amplitude due to both climate change and CO<sub>2</sub> increase (in the historical simulations) is nearly proportional to that due to CO<sub>2</sub> fertilization alone (from the 1%BGC runs), as shown by the best-fit straight line in Extended Data Fig. 3. As a result, the simulated increase in the CO<sub>2</sub> amplitude remains strongly correlated with the strength of the CO<sub>2</sub> fertilization across the model ensemble. This opens up the possibility of an emergent constraint<sup>8–11</sup> on CO<sub>2</sub> fertilization.

Figure 2 shows the extent of CO<sub>2</sub> fertilization in these same models at the time of doubling of CO<sub>2</sub> in the 1%BGC runs. The fractional increase in high-latitude (60° N–90° N) GPP due to the doubling of the CO<sub>2</sub> concentrations in these models varies from 20% to 60%, with four of the seven models giving values of less than 25% (Fig. 2b). There is a clear similarity to the histograms showing the sensitivity of the CO<sub>2</sub> amplitude at BRW and KMK to CO<sub>2</sub> (Fig. 1b, d).

The linear relationship between the CO<sub>2</sub> fertilization effect on the CO<sub>2</sub> amplitude at BRW and the relative GPP increase at the time of CO<sub>2</sub> doubling for the CMIP5 models is shown in Fig. 3a, with a correlation coefficient of  $r = 0.98$  ( $P = 0.0005$ ; for KMK  $r = 0.96$ ,  $P = 0.0004$ ). Models that simulate a large trend in the CO<sub>2</sub> amplitude at BRW also predict a large high-latitude GPP increase in the future. The combination of the observed trend in the CO<sub>2</sub> amplitude and this model-based linear relationship creates an emergent constraint<sup>8–11</sup> on the magnitude of CO<sub>2</sub> fertilization of large-scale high-latitude ecosystems in the real world. In the absence of this constraint, the prior probability density function (PDF) for the CMIP5 model spread is shown as the black histogram in Fig. 3b, implying a modal CO<sub>2</sub> fertilization effect of a 20%–25% increase in GPP due to a doubling of CO<sub>2</sub> at BRW. The observed range of the CO<sub>2</sub> fertilization effect at BRW allows a conditional PDF to be calculated by convolving the probability contours around the best-fit straight line in Fig. 3a with the uncertainty in the observed changes of the BRW CO<sub>2</sub> amplitude to annual mean CO<sub>2</sub> concentration<sup>10,11</sup>. This emergent constraint implies a reduced range of uncertainty for the CO<sub>2</sub> fertilization effect for high-latitude land, with a central estimate of  $37\% \pm 9\%$  that is also higher than those suggested by the unweighted CMIP5 models.

We use the same method for the KMK data–model comparison (Fig. 3c, d), but in this case we compare the CO<sub>2</sub> amplitude at KMK with the mean GPP in the entire extratropical Northern Hemisphere (30° N–90° N). We again find that the sensitivity of the CO<sub>2</sub> amplitude to the annual mean CO<sub>2</sub> has an approximately linear relationship with the magnitude of the CO<sub>2</sub> fertilization of GPP (Fig. 3c). This provides an emergent constraint on the CO<sub>2</sub> fertilization of extratropical GPP of  $32\% \pm 9\%$  due to doubling CO<sub>2</sub>, which overlaps with the estimate of CO<sub>2</sub> fertilization that we derived for BRW ( $37\% \pm 9\%$ ), providing further evidence of robust constraints.

For comparison, the Free Air CO<sub>2</sub> Enrichment (FACE) experiments suggest an increase in net primary productivity of about 23% when averaged across four sites with approximately 1.5 times the pre-industrial CO<sub>2</sub> concentration<sup>1</sup>, which is about 0.16% per p.p.m.v. Our larger-scale constraint therefore implies a similar central estimate of CO<sub>2</sub> fertilization on large scales and in the future, (approximately



**Figure 3 | Emergent constraints on the relative increase of large-scale GPP for a doubling of CO<sub>2</sub>.** **a, c,** Correlations between the sensitivity of the CO<sub>2</sub> amplitude to annual mean CO<sub>2</sub> increases at BRW (*x* axis) and the high-latitude (60° N–90° N) CO<sub>2</sub> fertilization on GPP at 2 × CO<sub>2</sub> (**a**) and the same for KMK and extratropical (30° N–90° N) GPP (**c**). The grey shading shows the range of the observed sensitivity. The red line shows the linear best fit across the CMIP5 ensemble together with the prediction error (orange) and error bars show the standard deviation for each data point. **b, d,** The probability density histogram for the unconstrained CO<sub>2</sub> fertilization of GPP (black) and the conditional PDF arising from the emergent constraints (red) for BRW (**b**) and KMK (**d**).

0.13% ± 0.03% per p.p.m.v. for high latitudes and 0.11% ± 0.03% per p.p.m.v. for the extratropics), but with significantly reduced uncertainties. Models without nitrogen limitations span the full range of CO<sub>2</sub> fertilization (20%–60%), whereas the current models that include nitrogen limitations appear to underestimate CO<sub>2</sub> fertilization (20%–25%), especially for the extratropical domain. These emergent constraints therefore give a consistent picture of a substantial CO<sub>2</sub>-fertilization effect and point to the need for further improvements in the treatment of nutrient limitations in ESMs.

**Online Content** Methods, along with any additional Extended Data display items and Source Data, are available in the online version of the paper; references unique to these sections appear only in the online paper.

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## METHODS

**Diagnosing the CO<sub>2</sub> fertilization factor.** The long-term CO<sub>2</sub> fertilization factor is defined in this study as the fractional change over time of GPP in the biogeochemically coupled simulation experiment (referred to as 1%BGC). The CO<sub>2</sub> fertilization factor was diagnosed individually for all models for the Northern Hemispheric high latitudes (60° N–90° N) and the extratropics (30° N–90° N) for a doubling of atmospheric CO<sub>2</sub> concentration from its pre-industrial value of 285 p.p.m.v. Individual values for each model are listed in Extended Data Tables 1 and 2. As not all models provided output from year zero, the fractional change was calculated from five-year means centred on year 10 and year 70 and divided by a factor of 0.9 to account for the missing first 10% of the CO<sub>2</sub> increase.

**Diagnosing the CO<sub>2</sub> effect on the CO<sub>2</sub> seasonal amplitude.** The amplitude of the seasonal cycle is derived from the difference between the maximum and minimum monthly mean atmospheric CO<sub>2</sub> concentrations for each year. To estimate the effect of increasing atmospheric CO<sub>2</sub> concentrations ( $\Delta C_A(t)$ , where  $C_A$  is the atmospheric CO<sub>2</sub> concentration and  $t$  is time) on the CO<sub>2</sub> seasonal cycle amplitude ( $\Delta C_{A,\text{ampl}}(t)$ ), we correlated these over the full length of the records available from the observations (see Methods section ‘Observational Data’) and the historical model simulations (1860–2005) for the CMIP5 models:

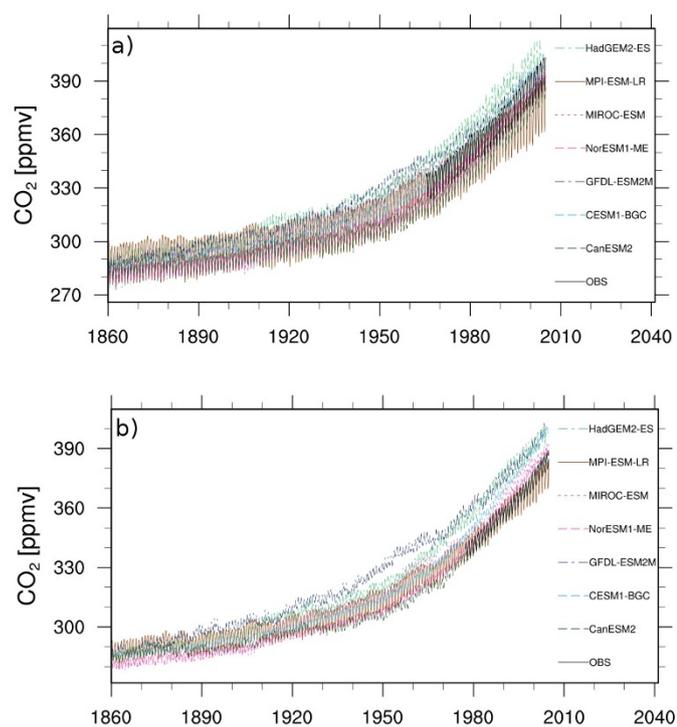
$$\Delta C_{A,\text{ampl}}(t) = a_0 + a \Delta C_A(t)$$

where  $a_0$  and  $a$  are fitting parameters. Individual values for each model resulting from this equation are given in Extended Data Tables 1 and 2.

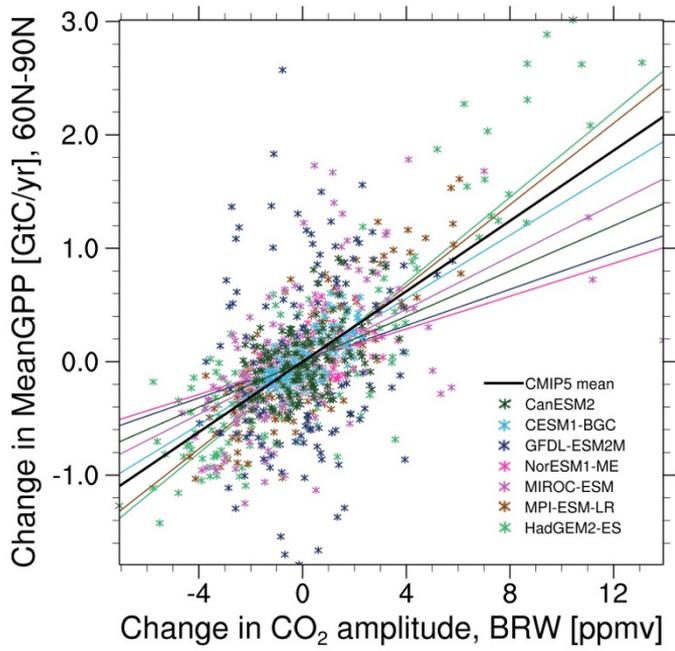
**Observational Data.** Observed monthly mean *in situ* atmospheric CO<sub>2</sub> concentrations at BRW (71.3° N, 156.6° W; record period 1974–2013) are from the National Oceanic and Atmospheric Administration (NOAA)/Earth System Research Laboratory (ESRL) (<http://www.esrl.noaa.gov/gmd/ccgg/trends>) and measurements at KMK (19.5° N, 155.6° W; record period 1979–present) from the Scripps Institute of Oceanography ([http://www.scrippsco2.ucsd.edu/research/atmospheric\\_co2](http://www.scrippsco2.ucsd.edu/research/atmospheric_co2)). Comparable data for the CMIP5 models were extracted as the near-surface CO<sub>2</sub> concentration for the closest grid box to each site.

**Code availability.** The routines used to reproduce this analysis from the CMIP5 model outputs are part of the ESMValTool<sup>19</sup> and are available upon request from the corresponding author.

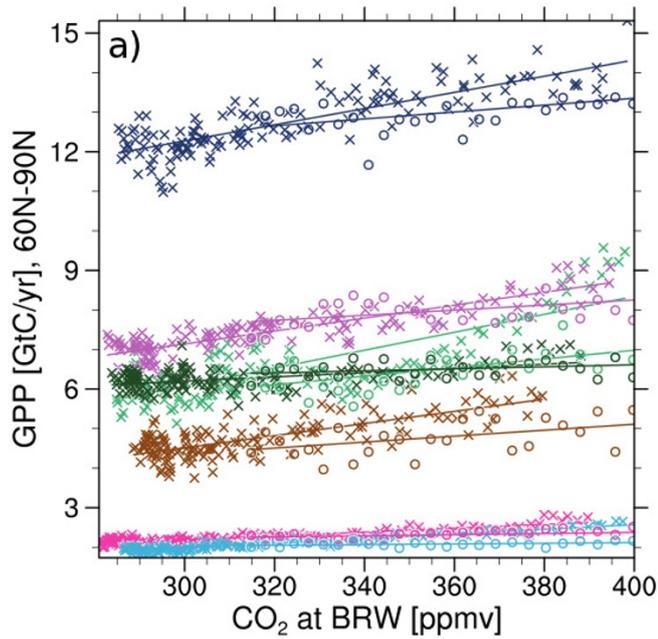
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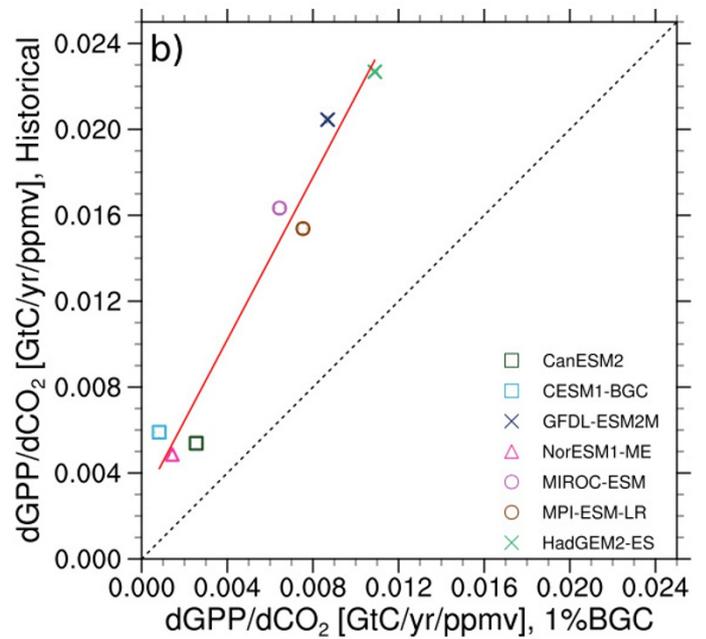
**Extended Data Figure 1 | Simulated and observed CO<sub>2</sub> concentrations.**  
**a, b,** Time series of monthly mean atmospheric CO<sub>2</sub> between 1860 and 2005 at BRW (**a**) and KMK (**b**) at surface level, as simulated by the CMIP5 models in the historical simulations and observed (black lines) at each measuring site.



**Extended Data Figure 2 | Annual mean high-latitude GPP (60° N–90° N) against the amplitude of the CO<sub>2</sub> seasonal cycle at BRW for each of the CMIP5 ESMs.** The markers show the values for the individual years between 1850 and 2005 for the CMIP5 historical simulations and lines show the linear best fit for each model. The black line indicates the multi-model mean of the CMIP5 models.



**Extended Data Figure 3 | Comparison of high-latitude GPP (60° N–90° N) versus annual mean CO<sub>2</sub>.** **a**, The correlation for both the historical (asterisks) and the 1%BGC (circles) CMIP5 model simulations. The markers show the values for the individual years and lines show the



linear best fit for each model. **b**, The comparison of the gradients in **a** for each model. The red solid line shows the gradient of the correlation and the black dashed line indicates a 1:1 correlation.

Extended Data Table 1 | Summary data for the changes in high-latitude GPP (60° N–90° N) and the amplitude of the CO<sub>2</sub> seasonal cycle at BRW

Model	GPP(2xCO <sub>2</sub> )/GPP(1xCO <sub>2</sub> )	Regression coefficients from linear regression	
		Offset to initial amplitude, a <sub>0</sub> [ppmv]	CO <sub>2</sub> sensitivity of amplitude, a [ppmv/ppmv]
CanESM2	1.18 ± 0.03	4.26 ± 1.10	0.020 ± 0.004
CESM1-BGC	1.22 ± 0.04	1.37 ± 1.26	0.023 ± 0.004
GFDL-ESM2M	1.25 ± 0.04	1.11 ± 1.62	0.027 ± 0.005
HadGEM2-ES	1.62 ± 0.11	-12.95 ± 1.53	0.108 ± 0.005
MIROC-ESM	1.38 ± 0.05	-2.36 ± 1.72	0.051 ± 0.005
MPI-ESM-LR	1.48 ± 0.12	2.18 ± 1.54	0.067 ± 0.005
NorESM1-ME	1.25 ± 0.04	2.86 ± 0.83	0.020 ± 0.003
<b>OBS</b>	-	<b>-2.54 ± 2.49</b>	<b>0.050 ± 0.007</b>

The fractional change in high-latitude GPP was calculated from the 1%BGC simulation for a doubling of CO<sub>2</sub> and is listed for each CMIP5 ESM. The regression coefficients result from a linear regression model for the change in the CO<sub>2</sub> amplitude at BRW (see Methods). Regressions are calculated using the full length of the records available from the CMIP5 historical simulations (1860–2005) and the BRW observations (1974–2013).

Extended Data Table 2 | Summary data for changes in extratropical GPP (30° N–90° N) and the amplitude of the CO<sub>2</sub> seasonal cycle at KMK

Model	GPP(2xCO <sub>2</sub> )/GPP(1xCO <sub>2</sub> )	Regression coefficients from linear regression	
		Offset to initial amplitude, a <sub>0</sub> [ppmv]	CO <sub>2</sub> sensitivity of amplitude, a [ppmv/ppmv]
CanESM2	1.33 ± 0.04	0.89 ± 0.63	0.010 ± 0.002
CESM1-BGC	1.26 ± 0.02	2.91 ± 0.31	0.000 ± 0.001
GFDL-ESM2M	1.44 ± 0.04	-1.07 ± 0.70	0.019 ± 0.002
HadGEM2-ES	1.59 ± 0.07	-6.65 ± 0.63	0.038 ± 0.002
MIROC-ESM	1.34 ± 0.04	-0.70 ± 0.57	0.019 ± 0.002
MPI-ESM-LR	1.56 ± 0.06	1.27 ± 0.90	0.031 ± 0.003
NorESM1-ME	1.21 ± 0.03	4.16 ± 0.35	-0.001 ± 0.001
<b>OBS</b>	-	<b>4.84 ± 0.85</b>	<b>0.008 ± 0.002</b>

The fractional change in extratropical GPP was calculated from the 1%BGC simulation for a doubling of CO<sub>2</sub> and is listed for each CMIP5 ESM. The regression coefficients result from a linear regression model for the change in the CO<sub>2</sub> amplitude at KMK (see Methods). Regressions are calculated using the full length of the records available from the CMIP5 historical simulations (1860–2005) and the KMK observations (1979–2015).